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Structural mapping using PALSAR data in the Central Gold Belt, Peninsular Malaysia



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REGEOLOGY REVIE

Amin Beiranvand Pour *, Mazlan Hashim

Institute of Geospatial Science & Technology (INSTeG), Universiti Teknologi Malaysia, 81310 UTM Skudai, Johor Bahru, Malaysia

A R T I C L E I N F O

ABSTRACT

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1. Introduction

The identification of geological structures and lineament analysis using remote sensing imagery are always considered complementary for any precious metals exploration program in arid and semi-arid regions (Abdelsalam et al., 2000; Kusky and Ramadan, 2002; Pour and Hashim, 2011, 2012a,b, 2013; Ramadan et al., 2001; Sabins, 1999). However, in tropical environments, the application of remote sensing data for geological structure mapping has been much more limited (Hashim et al., 2013), because of the persistent cloud coverage, limited bedrock exposures, and vegetation. Preliminary studies by Pour et al. (2013, 2014) and Pour and Hashim (2014) in the Bau gold mining district. Sarawak. East Malavsia. on the island of Borneo demonstrated the applicability of satellite remote sensing imagery for mineral exploration in tropical environments. These studies suggest that more investigation is required to test the application of remote sensing data for locating potential gold exploration targets in the Central Gold Belt (CGB) of the Peninsular Malaysia. Many gold mines and prospects in the Peninsular Malaysia are located in the Central Gold Belt (CGB) (Ariffin, 2012; Ariffin and Hewson, 2007; Scrivenor, 1931; Yeap, 1993). Gold mineralization in this belt is structurally controlled. The CGB is a highly potential region for prospecting gold exploration targets along the major lineament structures using remote sensing technology. To date, this gold belt has not been tested using recent generations of very high resolution satellite remote sensing imagery.

Synthetic Aperture Radar (SAR) is an active microwave remote sensing system which can acquire data regardless of day or night, cloud, haze or smoke over a region. SAR image data provide information different from that of optical sensors. Clouds are reasonably transparent to microwave providing measurements with almost any weather conditions. SAR images have been used for geological mapping in glaciated and vegetated terrain, structural geology investigations related to the search for mineral deposits and hydrocarbon traps, and studies of geologic hazards (Abdelsalam and Stern, 2000; Abdelsalam et al., 2000; Kusky and Ramadan, 2002; Pettinato et al., 2013; Pour and Hashim, 2014; Pour et al., 2013, 2014; Raharimahefa and Kusky, 2009; Ramadan et al., 2001; Singhroy, 1992; Thurmond et al., 2006; Zandbergen, 2008). Radar transmits and detects radiation with wavelengths between 2.0 and 100 cm, but typically at 2.5–3.8 cm (X-band), 4.0-7.5 cm (C-band), and 15.0-30.0 cm (L-band) (Campbell, 2007; Spatz, 1997; Woodhouse, 2006). Longer wavelengths optimize the depth of investigation of the radar signal and allow radar to have complete atmospheric transmission. Generally, the approximate depth of penetration is equal to radar's nominal wavelength. L-band can observe the forest's underlying surface features as well as the canopy because of its penetration capability (Abdelsalam et al., 2000; Henderson and Lewis, 1998; Shimada and Isoguchi, 2002). Thus, in tropical environments, L-band SAR data provide the possibility of obtaining more useable geological structure information from the ground than shorter wavelengths.

The Central Gold Belt (CGB) of Peninsular Malaysia has been investigated to map structural elements associated

with gold mineralization using the Phased Array type L-band Synthetic Aperture Radar (PALSAR) satellite remote

sensing data. Gold mineralization in this belt is structurally controlled and associated with steeply dipping faults

and fold hinges. Adaptive local sigma and directional filters were applied to PALSAR data for tracing structural

elements associated with gold mineralization. Structural features along the Bentong-Raub Suture Zone have been identified as highly potential areas for gold prospecting. Four sets of lineaments trending N-S, NE-SW,

NNW-SSE and ESE-WNW were identified. Results of this study demonstrate the applicability of PALSAR remote

sensing data to assist gold exploration in the CGB particularly in reducing costs related to exploration for

epithermal and polymetallic vein-type mineralization in tropical environments.

This research presents a remote sensing approach for geological structure mapping in tropical environments. The objectives of this study are (1) to map structural elements associated with gold mineralization in the Central Gold Belt of the Peninsular Malaysia using the

^{*} Corresponding author. Tel.: +60 7 5530666; fax: +60 7 5531174.

E-mail addresses: beiranvand.amin80@gmail.com, a.beiranvand@utm.my (A.B. Pour), mazlanhashim@utm.my, profmhashim@gmail.com (M. Hashim).

Phased Array type L-band Synthetic Aperture Radar (PALSAR) satellite remote sensing data at both regional and local scales and (2) to establish a cost-effective exploration approach for prospecting epithermal and polymetallic vein-type mineralization in tropical environments using PALSAR data.

2. Geological setting

Peninsular Malaysia forms an integral part of the Southeast Asian continental core of Sundaland and comprises two tectonic blocks/ terranes, the Sibumasu Terrane in the west and the Sukhothai Arc (East Malaya Block) in the east, which were assembled by the Late Triassic (Metcalfe, 2011, 2013a,b). Sibumasu and East Malaya are separated by the Bentong–Raub Suture Zone, which includes a tectonic mélange with ribbon-bedded cherts, schists, and minor ophiolites that represent Palaeo-Tethys remnants (Hutchison, 1975, 2009; Metcalfe, 2000). More than 90% of the plutonic rocks in the Peninsular Malaysia are granitic. The granitoids can be divided into two belts, a West Malaya Main Range S-Type group of granitoids that yield Late Triassic to earliest Jurassic, and an eastern Malaya group of dominantly I-Type granitoids with a range of ages from early Middle Permian to early Late Triassic (Searle et al., 2012; Sevastjanova et al., 2011).

Based on stratigraphical and structural differences the Peninsular Malaysia is divided into three geological belts: the Eastern Tin Belt, Central Gold Belt, and Western Tin Belt that are bounded by major fault zones (Hutchison, 1975; Khoo and Tan, 1983) (Fig. 1). The Central Belt of the Peninsular Malaysia is well-known as the Gold Belt (Scrivenor, 1931; Yeap, 1993). The Central Gold Belt (CGB) consists mainly of Permo-Triassic, low-grade metasediments, deep to shallow marine clastic sediments and limestone with abundant intermediate to acid volcanics and volcaniclastics, deposited in a paleo-arc basin (Metcalfe, 2002, 2011, 2013a,b). The belt coincides with the Bentong-Raub suture, which is a deep rooted 13 km wide north-south trending tectonic zone (Cocks et al., 2005; Tan, 1996). The CGB is bounded by the Bentong-Raub Suture Zone to the west and the Lebir Fault Zone to the East (Campi et al., 2002). The north-south trending Bentong-Raub Suture extends from Thailand through Raub and Bentong to the east of Malacca, Peninsular Malaysia (Fig. 1). This suture represents the main Palaeo-Tethys Ocean that was destroyed by collision between the Sibumasu and Sukhothai continental terranes of Southeast Asia (Metcalfe, 2000).

Major faults in the Peninsular Malaysia strike N–S, NNW–SSE, NW– SE, WNW–ESE, E–W, ENE–WSW and NE–SW and have undergone complex repeated movements, including microstructure evidence for both sinistral and dextral movements along many strike-slip faults (Shuib,



Fig. 1. Simplified geological map of the peninsular Malaysia. Modified from Metcalfe (2013a). Study area is located in black rectangle.

2009). Dating of faults in the Peninsular Malaysia is largely poorly constrained but major N–S trending faults are interpreted as the earliest and are related to oblique amalgamation of Sibumasu and the Sukhothai Arc in the Permian–Triassic. NNW–SSE major dextral faults are interpreted to be Late Triassic–Jurassic and to have resulted in the opening of the Jurassic–Cretaceous continental pull-apart basins. Reactivation of these faults as sinistral strike-slip faults occurred in the Late Cretaceous, synchronous with emplacement of granitoids, and deformation of Jurassic–Cretaceous red bed sequences (Shuib, 2009). Further reactivation of these faults occurred in the Cenozoic (Replumaz and Tapponnier, 2003).

Gold mineralization in the CGB is associated with the accretionary prism along the north–south trending terrain boundary of the Bentong–Raub Suture zone. Most of the gold mineralization took place within a low-grade meta-sedimentary-volcanic terrain formed during the collision of the Sibumasu block underneath the East Malaya (Sukhothai) block through the Permian to late Triassic (Ariffin and Hewson, 2007; Hutchison, 1996; Metcalfe, 2000). A collision structural overprint has generated major N–S and NW-SW trending left slip faults, and dilational Riedel and subsidiary shear zones and numerous splays associated with these faults (Hewson and Crips, 1992; Tjia and Zaitun, 1985).

The Bentong-Raub suture has accommodated considerable strikeslip movement. This has resulted in numerous splays running along the CGB (Metcalfe, 2000,b). The formation of the Bentong-Raub suture zone was probably coeval with the emplacement of major faults. The ore fluids ascended and deposited in structurally favorable traps, such as shear zones, saddle reefs, and fold hinges during metamorphism and deformation (Yeap, 1993). Major gold mineralization is observed along the steeply dipping faults and hosted in sandstone, carbonaceous shale, tuffaceous siltstone, and tuffaceous conglomerate (Ariffin and Hewson, 2007; Makoundi, 2012). These structures host many quartz-gold lodes within the CGB (Ariffin, 2012). Major primary gold mineralization patterns within the CGB can be grouped into two types (I and II). The type I mineralization consists of large quartz reefs/lodes and parallel swarms of veins that traverse metasediments and granite. Type II mineralization exhibits a broader variety of gold mineralization styles such disseminated gold with stockwork quartz veins affiliated with intrusive bodies, and volcanogenic exhalative sulfides within a shear zone system (Ariffin and Azizi, 1995; Pereira, 1993; Pereira et al., 1993). The ore mineral assemblages in both types include gold, pyrite, arsenopyrite, chalcopyrite, pyrrhotite, sphalerite, galena, and geochronite (Makoundi, 2012).

3. Materials and methods

3.1. PALSAR

Phased Array type L-band Synthetic Aperture Radar (PALSAR), onboard the Advanced Land Observing Satellite (ALOS), was launched on January 24, 2006 by an H-IIA rocket from Tanegashima Space Center. It was developed by Japanese Ministry of Economy, Trade and Industry (METI) as a joint effort with Japan Aerospace Exploration Agency (JAXA). PALSAR is an active microwave sensor for all-weather conditions observation and operable both day and night (ERSDAC, 2006; Igarashi, 2001; Rosenqvist et al., 2004). It has L-band synthetic aperture

Table	2
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Directional filters with 7*7 kernel matrix.

N–S						
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
-1.0000	-1.0000	-1.0000	0.0000	1.0000	1.0000	1.0000
F-W						
-1,0000	-1,0000	-10000	-1,0000	-1,0000	-1,0000	-10000
-1.0000	-1.0000	-1.0000	-1.0000	-1.0000	-1.0000	-1.0000
-1.0000	-1.0000	-1.0000	-1.0000	-1.0000	-1.0000	-1.0000
0.0000	0.0000	0.0000	-0.0000	-0.0000	-0.0000	-0.0000
1.0000	1.0000	1.0000	1.0000	1.0000	1.0000	1.0000
1.0000	1.0000	1.0000	1.0000	1.0000	1.0000	1.0000
1 0000	1 0000	1 0000	1 0000	1.0000	1.0000	1.0000
1.0000	1.0000	1.0000	1.0000			
1.0000 NE-SW	1.0000	1.0000	1.0000			
NE-SW - 1.4142	-1.4142	-1.4142	-0.7071	0.0000	0.0000	0.0000
NE-SW - 1.4142 - 1.4142	-1.4142 -1.4142	-1.4142 -1.4142	-0.7071 -0.7071	0.0000	0.0000	0.0000
NE-SW - 1.4142 - 1.4142 - 1.4142	-1.4142 -1.4142 -1.4142	-1.4142 -1.4142 -1.4142	-0.7071 -0.7071 -0.7071	0.0000 0.0000 0.0000	0.0000 0.0000 0.0000	0.0000 0.0000 0.0000
NE-SW - 1.4142 - 1.4142 - 1.4142 - 0.7071	-1.4142 -1.4142 -1.4142 -0.7071	-1.4142 -1.4142 -1.4142 -0.7071	-0.7071 -0.7071 -0.7071 0.0000	0.0000 0.0000 0.0000 0.7071	0.0000 0.0000 0.0000 0.7071	0.0000 0.0000 0.0000 0.7071
NE-SW - 1.4142 - 1.4142 - 0.7071 0.0000	-1.4142 -1.4142 -1.4142 -0.7071 0.0000	-1.4142 -1.4142 -1.4142 -0.7071 0.0000	-0.7071 -0.7071 -0.7071 0.0000 0.7071	0.0000 0.0000 0.0000 0.7071 1.4142	0.0000 0.0000 0.0000 0.7071 1.4142	0.0000 0.0000 0.0000 0.7071 1.4142
NE-SW - 1.4142 - 1.4142 - 1.4142 - 0.7071 0.0000 0.0000	$-1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000$	$-1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000$	$\begin{array}{c} -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \\ 0.7071 \\ 0.7071 \end{array}$	0.0000 0.0000 0.7071 1.4142 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142
NE-SW -1.4142 -1.4142 -0.7071 0.0000 0.0000 0.0000	$\begin{array}{r} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	$\begin{array}{r} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	$\begin{array}{c} -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \\ 0.7071 \\ 0.7071 \\ 0.7071 \end{array}$	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142	0.0000 0.0000 0.0000 0.7071 1.4142 1.4142 1.4142
NE-SW -1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000 0.0000 NW-SE	$\begin{array}{c} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	$\begin{array}{c} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	- 0.7071 - 0.7071 - 0.7071 0.0000 0.7071 0.7071 0.7071	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142
NE-SW -1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000 0.0000 NW-SE 0.0000	$\begin{array}{c} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	$\begin{array}{c} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	- 0.7071 - 0.7071 - 0.7071 0.0000 0.7071 0.7071 0.7071	0.0000 0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 - 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 - 1.4142	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 -1.4142
NE-SW - 1.4142 - 1.4142 - 0.7071 0.0000 0.0000 NW-SE 0.0000 0.0000	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ \end{array}$	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ \end{array}$	-0.7071 -0.7071 -0.7071 0.0000 0.7071 0.7071 0.7071 -0.7071 -0.7071	0.0000 0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 - 1.4142 - 1.4142	0.0000 0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 - 1.4142 - 1.4142	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ \end{array}$
NE-SW -1.4142 -1.4142 -1.4142 0.0000 0.0000 NW-SE 0.0000 0.0000 0.0000	$\begin{array}{c} -1.4142 \\ -1.4142 \\ -1.4142 \\ -0.7071 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \\ 0.0000 \end{array}$	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ \end{array}$	-0.7071 -0.7071 -0.7071 0.0000 0.7071 0.7071 -0.7071 -0.7071 -0.7071	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\end{array}$	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 -1.4142 -1.4142 -1.4142 -1.4142	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\end{array}$
NE-SW -1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000 NW-SE 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\end{array}$	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\end{array}$	$\begin{array}{c} -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \\ 0.7071 \\ 0.7071 \\ 0.7071 \\ -0.7071 \\ -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \end{array}$	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 - 1.4142 - 1.4142 - 1.4142 - 1.4142 - 0.7071	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 -1.4142 -1.4142 -1.4142 -1.4142 -0.7071	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\end{array}$
NOUL NE-SW -1.4142 -1.4142 -1.4142 -0.000 0.0000 NW-SE 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.7071 1.4142	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\end{array}$	-1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.7071 1.4142	-0.7071 -0.7071 -0.7071 0.7071 0.7071 0.7071 -0.7071 -0.7071 -0.7071 -0.7071 0.0000 0.7071	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000 \end{array}$	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000 \end{array}$	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000 \end{array}$
NUE-SW - 1.4142 - 1.4142 - 1.4142 - 0.7071 0.0000 0.0000 NW-SE 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 0.0000 1.4142 1.4142	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ \end{array}$	$\begin{array}{c} -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ \end{array}$	$\begin{array}{c} -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \\ 0.7071 \\ 0.7071 \\ 0.7071 \\ -0.7071 \\ -0.7071 \\ -0.7071 \\ 0.0000 \\ 0.7071 \\ 0.7071 \end{array}$	$\begin{array}{c} 0.0000\\ 0.0000\\ 0.0000\\ 0.7071\\ 1.4142\\ 1.4142\\ 1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -1.4142\\ -0.7071\\ 0.0000\\ 0.0000\\ \end{array}$	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 -1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000	0.0000 0.0000 0.7071 1.4142 1.4142 1.4142 -1.4142 -1.4142 -1.4142 -0.7071 0.0000 0.0000

radar with multi-mode observation function (Fine mode, Direct downlink, ScanSar mode, and Polarimetric mode) of multi-polarization configuration (HH, HV, VH, and VV), variable off-nadir angle (9.9–50.8°), switching spatial resolution (10, 30, and 100 m for Fine, Polarimetric, and ScanSar modes, respectively) and swath width observation (30, 70, and 250–350 km for Polarimetric, Fine and ScanSar modes, respectively) (ERSDAC, 2006; Igarashi, 2001).

PALSAR data can be used in specific fields, including (i) land area basin mapping (geological structural analysis of target areas), (ii) coastal area basin mapping, (iii) monitoring of environments and natural disasters, and (iv) research and development for the processing and application of multi-polarimetric SAR data (geological structural analysis during the first stage of resource exploration) (ERSDAC, 2006).

Generally, fine (high resolution) mode is the most frequently used observation mode with a ground resolution of up to 7 m, which enables detailed observation of the area of interest. Its maximum ground resolution of 7 m is one of the highest for a Synthetic Aperture Radar (SAR) loaded on a satellite. There are two kinds of observation modes, namely one is the observation mode by single polarization of HH or VV (FBS: high resolution mode, single polarization), and the other is the observation mode by dual polarization of HH + HV and VV + VH (FBD: high resolution mode, dual polarization). PALSAR can also simultaneously receive horizontal and vertical polarization per each polarized

Table 1

The characteristics of PALSAR Level 4.1 data used in this study (HH = horizontally transmitted and horizontally received, HV = horizontally transmitted and vertically received, VV = vertically transmitted and horizontally received).

Granule ID	Date of acquisation	Instrument operation mode	Polarity	Path/row
PASL4100706061550471312030001	2007/06/06	PLRM	HH + HV + VV + VH	485/8
PASL4100706061550391312030000	2007/06/06	PLRM	HH + HV + VV + VH	485/7
PASL4100709111556501312030004	2007/09/11	FBDH	HH + HV	488/7
PASL4100709111556581312030003	2007/09/11	FBDH	HH + HV	488/8

transmission, which is called multi-polarimetry. In addition, PALSAR can switch from horizontal to vertical polarization and vice versa at respective transmission pulse, enabling four polarizations by double simultaneous polarization, a function called full polarimetry (ERSDAC, 2006).

High resolution (fine mode), full polarimetry (multi-polarization mode), off nadir pointing function and other functions of PALSAR improved the accuracy of analyzing geological structure, distribution of rocks, and is expected to be used for the first stage of ore deposits exploration (ERSDAC, 2006). Consequently, PALSAR data are useful for geological structural analysis associated with epithermal or polymetallic vein-type mineralization especially in tropical regions, where optical sensors often fail due to bad weather conditions.

3.2. Data for the study area

In this investigation, two fine mode dual polarization and two polarimetric mode quad polarization level 4.1 PALSAR scenes were obtained from the Earth and Remote Sensing Data Analysis Center (ERSDAC)



Fig. 2. RGB color combination of PALSAR polarimetric HH, HV and VV images, southeastern part of the Bentong–Raub Suture Zone.

Japan (http://gds.palsar.ersdac.jspacesystems.or.jp/e/) for the Central Gold Belt (CGB) of the Peninsular Malaysia. The fine mode scenes used in this study contain high accuracy orbit data with good quality, 12.5 m pixel spacing, 16 bits per pixel, 83 km observation width in range direction, 81 km observation width in azimuth direction, incident angle 38.7°, and off-nadir angle of 34.3°. Polarimetric mode scenes have high accuracy orbit data with good quality, 25 m pixel spacing, 16 bits per pixel, 44 km observation width in range direction, 73 km observation width in azimuth direction, incident angle of 21.5°. Both datasets are geo-reference and geo-coded. It should be noted that the PALSAR images were acquired during the dry seasons. Table 1 shows the characteristics of PALSAR data used in this study.

3.3. Data pre-processing

Level 4.1 is the product of the high resolution mode by dual polarization or the product of polarimetry mode (or quad polarization). It is derived from processing Level 1.1 data with respect to dual polarization data of high resolution mode and quad polarization data of polarimetry mode. The processing includes (a) range compression using fast Fourier transform (FFT), (b) secondary range compression using range migration compensation, (c) range migration curvature corrections, (d) azimuth compression, (e) multi-look processing, and (f) conversion from coordination system from slant range to ground range (for only Geo-coded data). It consists of values for cross products (e.g. HH*HH, HH*HV) based on observed polarizations (HH, HV, VV and VH). Furthermore, these are the data with a slant range, which have calculated crossproducts capable of making Stoke's matrices to effectively utilize polarization data. The data are geo-reference and geo-coded data, which enables the images to be oriented so that the north direction of the observed image corresponds to the upper direction of the image. With geo-coded data, slant range is converted to ground range (Campbell, 2007; ERSDAC, 2006; Franceschetti and Lanari, 1999; Gelautz et al., 1998; Marino, 2012; Wise, 2002).

Radar images are inherently corrupted by speckle. The presence of speckle in an image reduces the detectability of ground targets, obscures the spatial patterns of surface features, and decreases the accuracy of automated image classification. Therefore, it is necessary to treat the speckle by filtering the data before it can be used in various applications (Lee and Jurkevich, 1994; Sheng and Xia, 1996; Sveinsson and Benediktsson, 1996).

To facilitate tracing the structural patterns and investigate the relationship between structural setting and gold mineralization in the Central Gold Belt (CGB) of the Peninsular Malaysia, Level 4.1 PALSAR data required to be filtered for speckle reduction. Adaptive filters remove radar speckle from images without seriously affecting the spatial



Fig. 3. Field photographs of the topographic expression of collision and compressional structures in Cameron Highlands.

characteristics of the data (Lopes et al., 1990; Research Systems, Inc., 2008; Shi and Fung, 1994). Therefore, adaptive filtering was applied to the Level 4.1 PALSAR data. The data were processed using the ENVI (Environment for Visualizing Images) version 4.8 software package.

In this study, the adaptive local sigma filter was selected and applied to accomplish speckle reduction and preserving both edges and features. The local sigma filter uses the local standard deviation computed for the filter box to determine valid pixels within the filter window. It replaces the pixel being filtered with the mean calculated using only the valid pixels within the filter box (Eliason and McEwen, 1990). 7*7 filter size in pixels was applied.

A default value of 1.000 was used for the sigma factor field. ENVI uses the sigma factor to determine which pixels are valid by calculating a minimum and maximum pixel value based on the number of standard deviations (sigma) entered and the local statistics. The pixel being filtered will be replaced by the average of surrounding valid pixels (Research Systems, Inc., 2008). Local sigma filter showed favorable output in preserving edges and features as well as speckle suppression in this study, and it seems to be more practical than other adaptive filters for geological structural analysis.

3.4. Data processing

Directional filtering technique was applied to the local sigma resultant image for detailed lineament extraction and edge enhancement. It is a spatial domain filtering technique and derivative edge enhancement filter that selectively enhances image features having specific direction components (gradients) (Carr, 1995; Haralick et al., 1987; Sabins, 1996; Vincent, 1997). Edge enhancement delineates the edge and makes the shapes and details comprising the image more conspicuous. It can be used in geological applications to highlight faults and lineaments. Directional filter is used for producing artificial effects suggesting tectonically controlled linear features (Amri et al., 2011; Drury, 1986; Kavak and Cetin, 2007; Suzen and Toprak, 1998). It is a straightforward method for extracting edges in the spatial domain that approximates the first derivative between two adjacent pixels. The algorithm produces the first difference of the image input in the horizontal, vertical, and diagonal directions (Jensen, 2005). As a result, many additional edges of diverse orientations are enhanced. Edge enhancement is performed by convolving the original data with a weighted mask or kernel. Chavez and Bauer (1982) suggested that the optimum kernel size (3*3, 5*5, 7*7, etc.) typically used in edge enhancement is a function of the surface roughness. Blurring becomes more severe as the size of the kernel increases, especially at the edges of objects (Jensen, 2005).

Directional filters were used to enhance specific linear trends in the local sigma resultant images. Four principal Directional filters: N–S, E–W, NE–SW, and NW–SE with a 7*7 kernel size were applied (Table 2). Filters were chosen to highlight the main lineament directions in the Central Gold Belt (CGB) of the Peninsular Malaysia. Directional filter angles were adjusted as N–S: 0°, E–W: 90°, NE–SW: 45°, and NW–SE: 135°. North (up) is zero degrees and the other angles are measured in the counterclockwise direction. 7*7 kernel matrix was selected to enhance semi-smooth and smooth/rough features. Image add back value was entered 60%. The image add back value is the percentage of the original image that is included in the final output image. This part of the original image preserves the spatial context and is typically done to sharpen an image.

3.5. Fieldwork

A global positioning system (GPS) survey was carried out using a Garmin® MONTANA® 650 to provide accurate locations for structural features associated with gold mineralization in the study area. Field view and outcrop photographs were taken of the geomorphology, hydrothermal alteration areas and structural elements. Additionally, image processing results were compared with the mineral distribution

map of the Peninsular Malaysia (1:500,000 scale) (Geological Survey Malaysia, 1988).

4. Results and discussion

4.1. Structural mapping of Bentong-Raub Suture Zone

PALSAR polarimetry observation data used in this study cover the southeastern part of Bentong–Raub Suture Zone. They were analyzed to illustrate the major lineaments trend and accentuate the tectonic structures of the region. Polarimetric mode observation function of PALSAR data has appropriate characteristics for structural mapping at a regional scale. It is possible to produce synthesized color images by allocating red–green–blue (RGB) color combination and placing them on each polarization data that are obtained by multi-polarization configuration (HH, HV, VH, and VV). RGB color-composite produces an image that depicts surface roughness associated with geological structures and lithology.



Fig. 4. RGB image of N–S (0°), NE–SW (45°), and NW–SE (135°) directional filters, southeastern part of the Bentong–Raub Suture Zone.

In this study, local sigma resultant images with different polarization configuration were assigned to RGB color-composite to provide visual interpretation of the Bentong-Raub Suture Zone. HH polarization image was assigned to red, HV polarization image was assigned to green, and VV polarization image was assigned to blue. Fig. 2 shows RGB color-composite image generated from PALSAR polarimetry observation data. Structural trends of the Bentong-Raub Suture Zone and collision and compressional structures in the Cameron Highlands are identified. Main orientations in the Bentong-Raub Suture Zone are N-S, NE-SW, and NW-SE. Water bodies appear black (south-western part of the image) and wet lands as mauve color (Fig. 2). Smooth surfaces such as calm water bodies appear dark in radar images due to reflection. The radar signal reflects away from the receiving antenna with an angle equal to that of the incident angle. In this case no returning radar signal will be detected in the receiving antenna (Abdelsalam and Stern, 2000; Thurmond et al., 2006). Fig. 3 shows a panoramic view of the topographic expression of collision and compressional structures in the Cameron Highlands.

For detailed mapping of lineament structures in the Bentong–Raub Suture Zone, directional filters were applied to HH, HV, VH, and VV polarization images. It seems that HV polarization is more suitable for lineament extraction and edge enhancement than other polarization images. Geological structures are more recognizable after directional filtering in the HV polarization image. Therefore, RGB color-composite was allocated to N–S, NE–SW, and NW–SE (R: 0°, G: 45°, B: 135°) with filtering directions derived from the HV polarization image (Fig. 4).

Two dominant directions can easily be identified, namely, N–S and NE–SW sets of lineaments (Fig. 4). More subtle lineaments strike approximately E–W and NW–SE. The continuous N–S striking lineament in the central to eastern part of the image corresponds to the boundary of the Bentong–Raub Suture Zone. N–S trending structures of the Bentong–Raub Suture Zone are apparent in Fig. 4. The collision zone and compressional structures appear clearly in the Cameron Highlands. Generally, most of the lineaments are clustered in the western part of the image. These lines mostly strike NNE.

PALSAR fine observation data were processed using Directional filters to map structural elements associated with known gold deposits in the Kuala Lipis region, Pahang and to identify areas along major lineaments in the Kelantan state that are prospective for ore deposits. Fine observation data have suitable spatial resolution (10 m) and a swath width (70 km), which enable detailed geological structural analysis of the study area at both regional and district scales.

4.2. Structural mapping of goldfields in Kuala Lipis, Pahang

PALSAR fine observation scene covering many of the gold mining districts in Kuala Lipis region in the state of Pahang was selected for detailed analysis of structural features associated with known gold



Fig. 5. District-scale geological map of the gold deposits, Kuala Lipis, Pahang, Central Malaysia. Modified from Makoundi (2012).



Fig. 6. RGB image of N–S (0°), NE–SW (45°), and NW–SE (135°) directional filters covering gold mining districts in Kuala Lipis region, Pahang.

deposits. It covers the eastern part of the Bentong–Raub Suture Zone and the central part of the CGB. Panjom (101° 58′ 58″ E, 4° 08′ 27″ N), Buffalo reef (101° 47′ 11″ E, 4° 15′ 59″ N), Selinsing (101° 47′ 38″ E, 4° 14′ 57″ N), Rubber hill (101° 47′ 55″ E, 4° 14′ 38″ N), Kechau-Tui (101° 58′ 49″ E, 4° 16′ 27″ N) and Tersang (101° 51′ 96″ E, 4° 04′ 81″ N) mines are located in this PALSAR scene. Fig. 5 shows the districtscale geological map of the selected study area.

Major gold mineralization is observed along the steeply dipping faults. Favorable settings for high-grade gold veins are the contact between tonalite and carbonaceous sedimentary rocks, especially where the latter are carbonaceous and/or strata are tightly folded or intensely faulted (Ariffin and Hewson, 2007). Fillis (2000) highlighted fold hinges as particularly well mineralized sites.

Directional filtering was implemented to the PALSAR data for tracing structural elements in the selected spatial scene covering the Panjom,



Fig. 8. Stratigraphic sequence of sedimentary rocks at Penjom ore deposit.

Buffalo reef, Selinsing, Rubber hill, Kechau-Tui and Tersang goldfields. Fig. 6 shows the RGB results for N–S, NE–SW, and NW–SE (R: 0°, G: 45°, B: 135°) filtering directions applied to the HV image. The above mentioned directional filters have been selected for RGB color application because NE-striking thrusts, NS-striking normal faults and NWstriking strike-slip faults are the most important structural elements for gold exploration in the CGB (Ariffin, 2012; Ariffin and Hewson, 2007).

Lineaments and form-lines are detected (Fig. 6), including the long lineaments and short lineaments that form linked systems with longer lineaments. The western and northern parts of the images exhibit longer and more lineaments than in the eastern part. Two major trends N and NE are mainly present in the western part of the image. The central and eastern parts of the image contain lineaments that strike NE and NW (Fig. 6). Lineaments mapped in the northern and central parts of the image express several fold systems as curvilinear structures. Lineaments associated with streams are interpreted to be fracture or fault controlled in the central north part.

N–S and NE–SW trending lineament systems are extensive in the region. Most longer lineaments strike N–S. N–S trending, normal-slip faults parallel to the Bentong–Raub Suture Zone trend are defined by a prominent west facing fault escarpment. This N–S trend is similar to the orientation of the Bentong–Raub Suture Zone (Fig. 6). Some NW trending lineaments are associated with normal faults. In the radar



Fig. 7. A regional view of the open-pit quarry of Penjom ore deposit.



Fig. 9. Association of hydrothermal alteration zones and lineament intersections with gold mineralization in Penjom ore deposit.



Fig. 10. Hydrothermally altered zone in Buffalo reef ore deposit.

image, strike-slip faults mark sharp boundaries; the planar fabric on either side is either sharply truncated or sheared (Abdelsalam et al., 2000).

Most of the known gold deposits are located along splay faults in the CGB, which are confined within brittle-ductile share or brecciated zones (Ariffin and Hewson, 2007; Yeap, 1993). Penjom gold deposit is located along splay faults. The Kelau–Karak fault (normal) is one of the major faults running across the Penjom goldfield that controls major plutonic emplacements (Tjia and Zaitun, 1985). Localized distribution of plutons is caused by faulting and folding, and the Penjom thrust has a NE–SW strike and southerly dip within the deposit (Ariffin, 2007). Major gold mineralization took place within the footwall of this thrust (Flindell, 2003). The Kelau–Karak fault and Penjom thrust are detected in the south-eastern part of the image (Fig. 6). A regional view of the open-pit quarry of Penjom is shown in Fig. 7. The Penjom thrust is the dominant feature controlling the distribution of ore at Penjom and strikes NE (035°) and dips to the southwest (30–40°) (Jasmi, 2007).

The stratigraphic sequence of sedimentary rocks at Penjom strikes N–S and dips moderately to the east (Fig. 8), which coincides with the regional N–S strike and with the trend of granitoid bodies. At Penjom goldfield, the lateral faults and shear zones are oriented in the N–S and NE–SW. Most high angle faults that strike N355°–005°E or N–S and N300°-310°E or WNW–ESE show right-lateral slip, whereas faults that strike N035°–045°E or NE–SW indicate left-lateral slip (Heru et al.,



Fig. 12. Association of hydrothermally altered zones and fault-related rocks with high grade ore mineralization section in the open pit quarry of Selinsing.

2000). Lineaments that are spatially associated with alteration zones are likely target areas for gold mineralization. Three dominant types of alteration were recognized in the Penjom gold deposit, including silicification, argillic alteration (illite) and chloritisation. Limonitic iron staining is also present (Wan and Heru, 2001, 2003). Fig. 9 shows the association between hydrothermal alteration zones, lineaments, and gold mineralization in the Penjom ore deposit.

The other major goldfields featured in Fig. 6 are the Buffalo reef, Selinsing, and Tersang. They are located along N–S trending regional structures in the NW and SW parts of the image (Fig. 6). N–S and NE–SW trending faults and fold systems are obviously manifested in the gold mining districts (Fig. 6).

The Buffalo reef lies close to the eastern flank of the Bentong–Raub Suture Zone. Gold mineralization in Buffalo reef is mainly confined to the marine clastic rock sequence, which mainly strikes in N-NE and dips south (Ariffin, 1995; Pereira, 1993). Significant gold mineralization is commonly hosted by an N–S trending sheared zone that cuts metamorphosed, brecciated and hydrothermally altered calcareous graphitic shale (Ariffin, 2012). Fig. 10 shows hydrothermally altered calcareous graphitic shale associated with gold mineralization in the Buffalo reef ore deposit.

Selinsing is located along the north striking Bentong–Raub Suture Zone, south of the Buffalo reef (Figs. 5 and 6). Fig. 11 shows the Selinsing open-pit. Gold mineralization areas are associated with N–S and NNE



Fig. 11. A regional view of the open-pit quarry of Selinsing ore deposit.



Fig. 13. A view of argillic alteration zone in the Tersang goldfield.

orientated faults and associated hydrothermal alteration zones (Mohd et al., 2009; Makoundi, 2012 (Fig. 12)).

Structural and textural investigations have shown two sets of NNW–SSE and NE–SW oriented mineralised veins at Tersang deposit (Fig.6). The argillic alteration in the Tersang deposit is characterized by sericite, illite, and montmorillonite (Fig. 13). Sericite alteration is located proximal to the high grade zone (Makoundi, 2012).

4.3. Prospecting potential areas in Kelantan state

PALSAR fine observation data scene that covers northern parts of the Bentong–Raub Suture Zone and the CGB were processed to map geological structures and identify prospective areas for gold mineralization along major lineaments in the Kelantan state. Gold mineralization is typically associated with hydrothermal quartz vein systems, skarns, and volcanogenic massive sulfides in the northern part of the CGB in Kelantan state. Structural element is one of the main controls on gold mineralization in this region (Ariffin, 2012).

Fig. 14 shows a selected spatial subset scene covering northern parts of the Bentong-Raub Suture Zone and CGB in the Kelantan state. RGB color combination was assigned for N-S (R: 0°), NE-SW (G: 45°), and NW-SE (B: 135°) directional filters. The HV polarization PALSAR image was selected to perform directional filtering. The N-S orientation of the Bentong-Raub Suture Zone is detected in the western part of the image (Fig. 14). The structural lines with NE-SW and NW-SE directions are in accordance with the tectonic framework of the study area. Major N-S and NE-SW orientations and strike-slip faults with sharp boundaries are revealed in Fig. 14. Several small faults and fractures with a NE-SW trend are clearly visible in the south and central part of the image. However, few small faults with N-S and NW-SE trends are observable in the north and eastern part of the image. Curvilinear features and some curving faults with NE-SW and NW-SE trends are apparent in the image especially in the central part (Fig. 14). These lineaments and curvilinear structures have great potential for hosting quartz-gold lodes.



Fig. 14. RGB image of N–S (0 $^{\circ}$), NE–SW (45 $^{\circ}$), and NW–SE (135 $^{\circ}$) directional filters covering the northern part of the CGB, Kelantan state.

Major gold deposits are located along the N–S Bentong–Raub Suture Zone in the CGB. Consequently, similar N–S trending structures are prospective for gold mineralization in the CGB. However, extensive deformation associated with the intersections of N–S, NE–SW, NNW– SSE and ESE–WNW brittle-ductile shear zones also represent favorable sites for gold mineralization at the district scale. Particularly favorable structural elements include the presence of fault-related rocks (cataclasite and mylonite) and coincident hydrothermal alteration. Curvilinear features are also important for gold prospecting. Fold hinges are favorable sites where they are associated with intensely faulted zones in contact between tonalite and carbonaceous sedimentary rocks. Consequently, the intersections of circular features, lineaments and hydrothermal alteration zones are important indicators for gold exploration in the CGB.

5. Conclusions

Results of this investigation provide an exploration approach using PALSAR data to map structural elements associated with gold mineralization along the Bentong–Raub Suture Zone in the Central Gold Belt (CGB) of the Peninsular Malaysia. Structural investigation has shown sets of N–S, NE–SW, NNW–SSE and ESE–WNW mineralized trends associated with fault-related rocks and hydrothermal alteration zones. These main fault trends are intersected by many shear or lateral fault zones. Of particular importance to exploration are shear zone, mylonite, cataclasite and felsic intrusive with coincident hydrothermal alteration. The results of this study demonstrate the usefulness of PALSAR satellite remote sensing data for mapping regional and district structural elements associated with epithermal and polymetallic vein-type mineralization in tropical environments.

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